

INVESTIGATION OF DRY AIR INTRUSION OVER THE ARABIAN SEA AND ITS IMPLICATIONS FOR INDIAN SUMMER MONSOON

RAHUL SINGH



CENTRE FOR ATMOSPHERIC SCIENCES

INDIAN INSTITUTE OF TECHNOLOGY DELHI

JULY 2024

© Indian Institute of Technology Delhi (IITD), New Delhi, 2024

**INVESTIGATION OF DRY AIR INTRUSION
OVER THE ARABIAN SEA AND ITS
IMPLICATIONS FOR INDIAN SUMMER
MONSOON**

by

RAHUL SINGH

CENTRE FOR ATMOSPHERIC SCIENCES

Submitted

in fulfillment of the requirements of the degree of

Doctor of Philosophy

to the



INDIAN INSTITUTE OF TECHNOLOGY DELHI

JULY 2024

Dedicated to ...

My beloved Parents

Mrs. Beena Singh &

Mr. Sanjay Singh

Certificate

This is to certify that the thesis entitled “**INVESTIGATION OF DRY AIR INTRUSION OVER THE ARABIAN SEA AND ITS IMPLICATIONS FOR INDIAN SUMMER MONSOON**”, submitted by **Mr. Rahul Singh** to the Indian Institute of Technology Delhi, for the award of the degree of **Doctor of Philosophy**, is a record of the original, bonafide research work carried out by him under my guidance and supervision. The thesis has reached the standards fulfilling the requirements of the regulations related to the award of the degree.

The results contained in this thesis have not been submitted in part or in full to any other University or Institute for the award of any degree or diploma.

Date:

Place: New Delhi

Prof. Sandeep Sukumaran

Associate Professor

Centre for Atmospheric Sciences

Indian Institute of Technology Delhi

New Delhi-110016, India

Acknowledgements

I would like to begin by expressing heartfelt gratitude to God, whose divine grace has illuminated my path and provided me with strength and perseverance throughout this journey. I am profoundly thankful for the blessings, guidance, and protection bestowed upon me, enabling me to navigate the challenges and triumphs encountered along the way.

I would also like to extend my heartfelt thanks and recognition to all the individuals and organizations who have provided their invaluable support and encouragement throughout the completion of this thesis, as this work would not have been feasible without their assistance.

*I would like to express my sincere gratitude to my supervisor, **Prof. Sandeep Sukumaran**, for his invaluable guidance, support, and encouragement throughout this research journey. His expertise, patience, and unwavering commitment to excellence have been crucial in shaping this thesis. I am grateful for the opportunities he has provided, his support during challenging times, and his encouragement to strive for excellence. His mentorship has not only improved my skills and knowledge but also inspired me to achieve new heights in my pursuits. I am grateful for his dedication to my success and privileged to work under his exemplary guidance.*

*I am also grateful to the members of my Student Research Committee (SRC) - **Prof. Krishna AchutaRao**, **Prof. Sarvesh Kumar Dubey**, and **Prof. Vamsi Chalamalla** for their valuable feedback and constructive criticism, which have greatly improved the quality of this work. I would also like to express my appreciation to **Prof.***

Somnath Baidya Roy, the Head of the Centre for Atmospheric Sciences (CAS), for providing me with a conducive workspace and necessary facilities.

My heartfelt thanks go to the Indian Institute of Technology Delhi, India, for providing the resources and facilities necessary for carrying out this research.

I am deeply grateful to the most important people in my life for their love, understanding, and unwavering belief in my abilities: my dear parents, **Mrs. Beena Singh** and **Mr. Sanjay Singh**, and last but not least, my younger brother, **Rohit Singh**. Their consistent encouragement and support have been my main sources of strength and motivation.

I would like to express my gratitude to my friends and colleagues for their support, guidance, and companionship during this demanding yet fulfilling experience. Their presence has made difficult times easier and happy moments more special. Thank you, dear friends, for always being there, encouraging me, and bringing joy and laughter into my life.

This thesis would not have been possible without the collective contributions and support of all those mentioned above. Thank you for being a part of this transformative experience.

Abstract

The dry air intrusion over India during summer monsoon break phases is well known. It has been argued that this dry air originates over the desert regions of West Asia. During the summer monsoon season, a reservoir of saturation deficit air is identified over the western and northern Arabian Sea in the present study. The monsoon low-level jet (LLJ) which transports the moisture to continental India in the active phase of monsoon, transports the dry air to northern and central India during the break phase. The LLJ undergoes a weakening and broadening before the monsoon break phase in response to increased barotropic instability. The broadening of the LLJ leads to an intensification of zonal flow in the poleward flanks and a weakening at the core. The development of a positive meridional gradient in sea surface temperature over the northern Arabian Sea favors an increase in the low-level zonal flow in the north, which advects the moist deficit air across northwest India. The dry air intrusion results in enhanced static stability over northern and central India and a strong suppression of convection. Further, the enhanced static stability weakens zonal flow from the northern Arabian Sea region and leads to the demise of the dry air intrusion. Thus, internal mechanisms are responsible for the dry air intrusion over India and its termination during the break phase of the summer monsoon. An index for the dry air intrusion is constructed based on the saturation deficit transport. This dry air intrusion index is used to identify the dry air intrusion events during monsoon breaks. The statistics show that most of the break events are found to be in close coincidence with dry air intrusion activity. It is also observed that the dry air intrusion and the monsoon breaks are happening simultaneously, suggesting that it is difficult to establish a cause-effect relationship.

It is now evident that the strengthening of monsoon LLJ in the northern parts of the Arabian Sea during the break phase of Indian summer monsoon (ISM) helps in transporting this dry air towards northwestern India. Hence, over northwestern India (NWI), a weakening (strengthening) of the zonal flow over the northern Arabian Sea can reduce (enhance) the influx of unsaturated air to the NWI and thereby enhance (reduce) precipitation there. The variability in the zonal flow over the northern Arabian Sea is a direct geostrophic response to the variability in the meridional pressure gradient over NWI. The inter-annual variability in the mean sea

level pressure over the region explains the inter-annual variability of ISM precipitation during July-August over northwestern India. The contribution of El Niño Southern Oscillation in the inter-annual variability of precipitation over this region is not significant.

Although the existence of dry air intrusion is well-documented through observations, its representation in climate models remains uncertain. It is important to enhance our understanding of the process of dry air advection in climate models to assess their fidelity in simulating the climate over the region. Most climate models participated in the sixth phase of Coupled Model Intercomparison Project (CMIP6) models analyzed in this study realistically simulate the observed pattern of dry air advection over continental India during the summer monsoon break phase. Some models also simulate dry air transport from West Asia, possibly due to an overly smoothed representation of orography. Furthermore, the majority of CMIP6 models successfully capture the intrinsic modes associated with the dry monsoon phase, as demonstrated by empirical orthogonal function analysis on low-level zonal winds. It suggests that the global climate models exhibit better skill in simulating the dry processes of the monsoon compared to moist processes. These findings uncover previously under-explored aspects of the monsoon, which are essential for accurately assessing future regional climate changes. Further, the orography of the Hindu Kush region significantly influences precipitation variability in continental India by controlling the dynamics of dry air advection. Therefore, accurate representation of orography in climate models is essential to reduce precipitation bias. Many studies have reported a poleward shift of the monsoon LLJ in recent decades, which is projected to continue in a warming climate. The role of the LLJ in a warming climate is crucial, as it is the primary carrier of dry air to continental India. In a warming climate, the LLJ becomes more zonal in some models, leading to an increase in break days in the SSP 585 scenario, and becomes inclined, parallel to the Arabian Peninsula coastline, when break days decrease. This change may be due to the strengthening or weakening of the meridional component of horizontal wind, which affects moisture transport from the Arabian Peninsula and the northern Arabian Sea.

सार

ग्रीष्म मानसून विराम के चरणों के दौरान भारत में शुष्क हवा का आक्रमण सर्वविदित है। यह तर्क दिया गया है कि यह शुष्क हवा पश्चिम एशिया के रेगिस्तानी क्षेत्रों से उत्पन्न होती है। ग्रीष्मकालीन मानसून के मौसम के दौरान, वर्तमान अध्ययन में पश्चिमी और उत्तरी अरब सागर के ऊपर संतृप्ति की कमी वाली हवा के भंडार की पहचान की गई है। मानसून निम्न-स्तरीय जेट (एलएलजे) जो मानसून के सक्रिय चरण में नमी को महाद्वीपीय भारत तक पहुंचाता है, ब्रेक चरण के दौरान शुष्क हवा को उत्तरी और मध्य भारत में पहुंचाता है। बढ़ी हुई बारोट्रोपिक अस्थिरता के जवाब में मानसून ब्रेक चरण से पहले एलएलजे कमजोर और चौड़ा हो जाता है। एलएलजे के चौड़ीकरण से ध्रुवीय किनारों में आंचलिक प्रवाह तेज हो जाता है और कोर कमजोर हो जाता है। उत्तरी अरब सागर के ऊपर समुद्र की सतह के तापमान में एक सकारात्मक मेरिडियनल ग्रेडिएंट का विकास उत्तर में निम्न-स्तरीय आंचलिक प्रवाह में वृद्धि का पक्ष लेता है, जो पूरे उत्तर पश्चिम भारत में नम हवा की कमी को दर्शाता है। शुष्क वायु घुसपैठ के परिणामस्वरूप उत्तरी और मध्य भारत में स्थैतिक स्थिरता बढ़ जाती है और संवहन का एक मजबूत दमन होता है। इसके अलावा, बढ़ी हुई स्थैतिक स्थिरता उत्तरी अरब सागर क्षेत्र से क्षेत्रीय प्रवाह को कमजोर कर देती है और शुष्क हवा की घुसपैठ को समाप्त कर देती है। इस प्रकार, आंतरिक तंत्र भारत में शुष्क हवा की घुसपैठ और ग्रीष्मकालीन मानसून के ब्रेक चरण के दौरान इसकी समाप्ति के लिए जिम्मेदार हैं। संतृप्ति घाटे के परिवहन के आधार पर शुष्क वायु घुसपैठ के लिए एक सूचकांक का निर्माण किया जाता है। इस शुष्क वायु घुसपैठ सूचकांक का उपयोग मानसून अवकाश के दौरान शुष्क वायु घुसपैठ की घटनाओं की पहचान करने के लिए किया जाता है। आँकड़े बताते हैं कि ब्रेक की अधिकांश घटनाएँ शुष्क वायु घुसपैठ गतिविधि के साथ घनिष्ठ संयोग में पाई जाती हैं। यह भी देखा गया है कि शुष्क हवा का प्रवेश और मानसून का टूटना एक साथ हो रहा है, जिससे पता चलता है कि कारण-प्रभाव संबंध स्थापित करना मुश्किल है।

अब यह स्पष्ट है कि भारतीय ग्रीष्मकालीन मानसून (आईएसएम) के ब्रेक चरण के दौरान अरब सागर के उत्तरी हिस्सों में मानसून एलएलजे की मजबूती इस शुष्क हवा को उत्तर-पश्चिमी भारत की ओर ले जाने में मदद करती है। इसलिए, उत्तर-पश्चिमी भारत (एनडब्ल्यूआई) में, उत्तरी अरब सागर के ऊपर आंचलिक प्रवाह के कमजोर (मजबूत होने) से एनडब्ल्यूआई में असंतृप्त हवा का प्रवाह कम (बढ़) सकता है और इस तरह वहां वर्षा में वृद्धि (कम) हो सकती है। उत्तरी अरब सागर के ऊपर आंचलिक प्रवाह में परिवर्तनशीलता एनडब्ल्यूआई पर मध्याह्निय दबाव प्रवणता में परिवर्तनशीलता के लिए एक सीधी भूगर्भीय प्रतिक्रिया है। क्षेत्र पर औसत समुद्री स्तर के दबाव में अंतर-वार्षिक परिवर्तनशीलता उत्तर-पश्चिमी भारत में जुलाई-अगस्त के दौरान आईएसएम वर्षा की अंतर-वार्षिक परिवर्तनशीलता को बताती है। इस क्षेत्र में वर्षा की अंतर-वार्षिक परिवर्तनशीलता में अल नीनो दक्षिणी दोलन का योगदान महत्वपूर्ण नहीं है।

यद्यपि शुष्क वायु घुसपैठ का अस्तित्व अवलोकनों के माध्यम से अच्छी तरह से प्रलेखित है, जलवायु मॉडल में इसका प्रतिनिधित्व अनिश्चित बना हुआ है। क्षेत्र में जलवायु का अनुकरण करने में उनकी निष्ठा का आकलन करने के लिए जलवायु मॉडल में शुष्क वायु संवहन की प्रक्रिया के बारे में हमारी समझ को बढ़ाना महत्वपूर्ण है। अधिकांश जलवायु मॉडलों ने युग्मित मॉडल इंटरकंपेरिसन प्रोजेक्ट (सीएमआईपी6) के छठे चरण में भाग लिया, इस अध्ययन में विश्लेषण

किए गए मॉडल ग्रीष्मकालीन मानसून ब्रेक चरण के दौरान महाद्वीपीय भारत पर शुष्क वायु संवहन के देखे गए पैटर्न का वास्तविक रूप से अनुकरण करते हैं। कुछ मॉडल पश्चिम एशिया से शुष्क वायु परिवहन का भी अनुकरण करते हैं, संभवतः ऑरोग्राफी के अत्यधिक सुचारु प्रतिनिधित्व के कारण। इसके अलावा, अधिकांश सीएमआईपी6 मॉडल शुष्क मानसून चरण से जुड़े आंतरिक तरीकों को सफलतापूर्वक पकड़ लेते हैं, जैसा कि निम्न-स्तरीय क्षेत्रीय हवाओं पर अनुभवजन्य ऑर्थोगोनल फ़ंक्शन विश्लेषण द्वारा प्रदर्शित किया गया है। यह सुझाव देता है कि वैश्विक जलवायु मॉडल नम प्रक्रियाओं की तुलना में मानसून की शुष्क प्रक्रियाओं का अनुकरण करने में बेहतर कौशल प्रदर्शित करते हैं। ये निष्कर्ष मानसून के पहले से कम खोजे गए पहलुओं को उजागर करते हैं, जो भविष्य के क्षेत्रीय जलवायु परिवर्तनों का सटीक आकलन करने के लिए आवश्यक हैं। इसके अलावा, हिंदू कुश क्षेत्र की भौगोलिक स्थिति शुष्क वायु संवहन की गतिशीलता को नियंत्रित करके महाद्वीपीय भारत में वर्षा परिवर्तनशीलता को महत्वपूर्ण रूप से प्रभावित करती है। इसलिए, वर्षा पूर्वाग्रह को कम करने के लिए जलवायु मॉडल में भौगोलिक स्थिति का सटीक प्रतिनिधित्व आवश्यक है। कई अध्ययनों ने हाल के दशकों में मानसून एलएलजे में ध्रुवीय बदलाव की सूचना दी है, जिसके गर्म जलवायु में जारी रहने का अनुमान है। गर्म होती जलवायु में एलएलजे की भूमिका महत्वपूर्ण है, क्योंकि यह महाद्वीपीय भारत के लिए शुष्क हवा का प्राथमिक वाहक है। गर्म जलवायु में, एलएलजे कुछ मॉडलों में अधिक क्षेत्रीय हो जाता है, जिससे एसएसपी 585 परिदृश्य में ब्रेक के दिनों में वृद्धि होती है, और जब ब्रेक के दिन कम हो जाते हैं, तो यह अरब प्रायद्वीप समुद्र तट के समानांतर झुका हुआ हो जाता है। यह परिवर्तन क्षैतिज हवा के मध्याह्न घटक के मजबूत होने या कमजोर होने के कारण हो सकता है, जो अरब प्रायद्वीप और उत्तरी अरब सागर से नमी परिवहन को प्रभावित करता है।

Contents

Certificate	i
Acknowledgements	iii
Abstract	v
Contents	ix
List of Figures	xiii
List of Tables	xxiii
Abbreviations	xxv
1 Introduction	1
1.1 Background	1
1.1.1 Indian Summer Monsoon	1
1.1.2 Analyzing the factors contributing to prolonged drought conditions over the continental India	5
1.2 Motivation and Research Gaps	9
1.3 Objectives of the thesis	11
1.4 Structure of the thesis	12
2 Data and Methods	15
2.1 Observation data	15
2.1.1 Precipitation data	15
2.1.1.1 IMD data	15
2.1.1.2 GPCP data	16
2.1.1.3 TRMM data	16
2.1.2 OLR data	17
2.1.3 Sea Surface Temperature data	17
2.1.3.1 Daily SST data	17
2.1.3.2 Monthly SST data	17

2.2	Reanalysis data	18
2.2.1	ERA-Interim	18
2.2.2	ERA5	19
2.3	Climate model data	19
2.4	Methodology	20
2.4.1	Identification of active and break spells	20
2.4.2	Identification of dry air intrusion events	20
2.4.3	EOF analysis	21
2.4.4	Isentropic analysis	22
2.4.5	Atmospheric General Circulation Model Simulations	22
2.4.6	Statistical significance test	23
3	Dynamics of dry air intrusion over India during summer monsoon breaks	25
3.1	Introduction	25
3.2	Datasets and Methodology	26
3.2.1	Datasets overview	26
3.2.2	Methods	26
3.3	Results and Discussion	27
3.3.1	Climatological analysis	27
3.3.2	EOF and Regression analysis	29
3.3.3	Characterisation of break events and its associated dry air advection activity.	32
3.3.4	Thermodynamic stability of the atmosphere	38
3.4	Summary and Conclusions	46
4	A Mechanism for the Summer Monsoon Precipitation Variability over Northwest India Driven by Moisture Deficit Transport	49
4.1	Introduction	49
4.2	Datasets and Methodology	49
4.2.1	Datasets overview	50
4.2.2	Methods	51
4.3	Results and Discussion	53
4.3.1	Climatology of precipitation over Northwest India	53
4.3.2	Dynamics associated with mean sea level pressure and dry air advection	54
4.3.3	The role of ENSO and Silk Road Pattern	62
4.3.4	Lead-lag relationship between precipitation and dry air advection	63
4.4	Summary and Conclusions	69

5	Investigation of dry air intrusion activity over India during the summer monsoon in climate model simulations	73
5.1	Introduction	73
5.2	Datasets and Methodology	74
5.2.1	Datasets overview	74
5.2.2	Methods	75
5.3	Results and Discussion	79
5.3.1	Dry air intrusion over India during break phases of the Indian summer monsoon in CMIP6 models	79
5.3.2	The Role of Orography in Simulating Dry Air Advection over India during the Summer Monsoon	101
5.3.3	LLJ dynamics over the Arabian Sea during the Indian summer monsoon and its interaction with dry air intrusion activity in a warming climate	104
5.4	Summary and Conclusions	113
6	Conclusions and Scope of Future Work	117
6.1	Conclusions	117
6.2	Scope of Future Work	119
	References	121
	List of Publications	137
	Curriculum Vitae	139

List of Figures

1.1	Seasonal mean climatology of precipitation (mm day^{-1}). The analysis period is from 1998 to 2019, during June to September (JJAS) months. TRMM [<i>Huffman et al.</i> , 2016] dataset is used.	2
1.2	Ten-day back trajectories from central India. (left) Dry spell for trajectories terminating at 850 (green) and 700 (red) hPa; (right) as at left, but for the wet spells of the Indian summer monsoon, for (a) 18 Jun 2009, (b) 14 Aug 2005, (c) 16 Jul 2002, (d) 30 Aug 2001, (e) 14 Jul 2009, (f) 1 Aug 2005, (g) 1 Aug 2004, and (h) 12 Jul 2001 (<i>Krishnamurti et al.</i> [2010]; Published 2010 by the American Meteorological Society).	8
2.1	Schematic of the calculation of dry air intrusion index (DAI). DAI is computed by standardizing the daily vertically integrated anomalies of zonal moisture deficit transport from 950 hPa to 300 hPa averaged over $20^{\circ}\text{N} - 30^{\circ}\text{N}$ at 70°E	21
2.2	Region chosen to see the effect of dry air intrusion on Indian Summer Monsoon by changing orography in CAM5 experiments.	23
3.1	Climatology of saturation deficit (contours; g kg^{-1}) at 850hPa and precipitation (shading; mm day^{-1}) for the months of (a) July and (b) August. The climatology is calculated for the period of 1998 – 2017 using TRMM daily precipitation	28
3.2	(a) First and (b) second leading modes of EOF of daily zonal wind anomalies at 850 hPa for the year 2009; (c) Standardized PC time series corresponding to the first two leading modes of EOF. The sign of EOF2 and the corresponding PC time series is flipped.	30
3.3	Regression of saturation deficit transport vectors ($\mathbf{V}_{\text{qd}}$) on PC2 of daily zonal wind anomalies at 850 hPa. The shading shows statistically significant ($p < 0.05$) regression coefficients of magnitude of saturation deficit transport on PC2. Statistical significance is estimated using a two-tailed t -test. PC2 is multiplied by -1 before regression. The calculation is done for July-August period during 1981-2014. . . .	31
3.4	Time series of dry air intrusion index, standardized daily precipitation anomalies over central India, and standardized principal component time series corresponding to the second leading mode of the EOF of zonal wind anomalies at 850 hPa. A three-day running mean is applied to all the time series. All calculations are done for July - August, 2009	34

3.5	Composites of saturation deficit (shading; g kg^{-1}) and wind vectors (m s^{-1}) at 850 hPa for (a) break spells without dry air intrusion and (b) dry air intrusion days during break spells. The composite (solid line) and ± 1 standard deviation (dashed lines) of zonal mean ($55^\circ\text{E} - 75^\circ\text{E}$) zonal winds are shown as a function of latitudes. The red line in Fig. 3.5a shows the latitude band over which the dry air intrusion index is calculated. Only the statistically significant ($p < 0.05$) composites of saturation deficit are shown.	35
3.6	Lead-lag composite of daily SST ($^\circ\text{C}$) for (a) 4 days before, (b) 2 days before, (c) day of, and (d) 2 days after the onset of dry air intrusion. The composites are constructed for the 34 dry air intrusion events during July-August 1981 - 2014.	36
3.7	Lead-lag composite of area averaged zonal winds (black line) and d^2u/dy^2 (blue line) for the dry air intrusion events. The lag-0 corresponds to the first day of intrusion. The area average is done over $55^\circ\text{E} - 75^\circ\text{E}$ and $10^\circ\text{N} - 20^\circ\text{N}$	37
3.8	Composites of time evolution of vertical structure of equivalent potential temperature anomalies (K) averaged over (a) northwestern India ($65^\circ\text{E} - 75^\circ\text{E}$, $20^\circ\text{N} - 30^\circ\text{N}$) and (b) central India ($75^\circ\text{E} - 85^\circ\text{E}$, $15^\circ\text{N} - 25^\circ\text{N}$) for dry air intrusion events (shading) and for monsoon breaks without dry air intrusion (contour). Lag 0 corresponds to the day of maximum of dry air intrusion. In the case of no intrusion composite, lag 0 corresponds to the median of the monsoon break period. The dry air intrusion composites are constructed for 34 events during July-August 1981 - 2014 while the no intrusion composites are constructed for four events in the same period.	39
3.9	Lead-lag composites of daily zonal wind anomalies (m s^{-1}) at $\theta = 316\text{K}$ isentropic level for all dry air intrusion events in July-August during 1981 - 2014. t_0 corresponds to the day of maximum of dry air intrusion. The stippling shows the region where the composite is statistically significant ($p < 0.05$) as revealed by a two-tailed t test.	41
3.10	Same as Fig. 3.9 except for daily specific humidity anomalies (g kg^{-1}). The stippling shows the region where the composite is statistically significant ($p < 0.05$) as revealed by a two-tailed t test.	42
3.11	Same as Fig. 3.9 except for the daily anomalies of relative vorticity (shading; $\times 10^{-5}\text{s}^{-1}$) and wind vectors (m s^{-1}). The stippling shows the region where the composite of relative vorticity anomaly is statistically significant ($p < 0.05$) as revealed by a two-tailed t test.	43
3.12	Lead-lag composite of daily equivalent potential temperature anomalies (K) at 700hPa, with respect to the date of maximum of dry air intrusion (t_0). The stippling shows the region where the composite is statistically significant ($p < 0.05$) as revealed by a two-tailed t test.	44

3.13	Lead-lag composites of daily anomalies of outgoing longwave radiation (W m^{-2}) with respect to the day of maximum dry air intrusion (t_0). The stippling shows the region where the composite is statistically significant ($p < 0.05$) as revealed by a two-tailed t test.	45
4.1	(a) July-August climatology and composite anomalies of rainfall (mm day^{-1}) for (b) excess rainfall years and (c) deficit rainfall years. July-August months are chosen to compute the seasonal mean for each year, and the green rectangular box shows the area chosen over the northwestern region for the analysis. The excess (deficit) years are defined as years in which standardized anomaly of July-August seasonal mean rainfall ≥ 1 (≤ -1). The climatology is computed for 1951-2019.	50
4.2	Spatial linear trend map of (a) OLR ($\text{W m}^{-2} \text{ year}^{-1}$) and (b) Precipitation ($\text{mm day}^{-1} \text{ year}^{-1}$) for seasonal mean analysis. June to September (JJAS) months are chosen to compute the seasonal mean for each year. The analysis is performed for the period 1981-2022. The green box is the area chosen over the northwestern India for the analysis. The dashed lines show regions with statistically significant ($p < 0.05$) trend as revealed by two-tailed t -test.	52
4.3	(a) Climatology and composite difference analysis of geostrophic wind vectors at 850-hPa (ms^{-1} ; vectors) and mean sea level pressure (hPa; shading) (b) for the excess rainfall years when seasonal mean standardized rainfall anomaly ≥ 1 and (c) for the deficit rainfall years when seasonal mean standardized rainfall anomaly ≤ -1 . Difference is computed by subtracting long term climatology (1951-2019; a) from each of the above defined cases (b, c). July-August months are chosen to compute the seasonal mean for each year, and the gray rectangular box (22°N - 28°N , 62°E - 68°E) shows the area chosen over the northern Arabian Sea region for the regression analysis (Fig. 4.9). The area with gray shading represents the region where the climatology (1951-2019) of surface pressure is less than 850 hPa.	55
4.4	Composite difference analysis of moisture deficit (g kg^{-1} ; shading), wind vectors (ms^{-1} ; vectors) and temperature (K; shading) at 850-hPa. (a, c) for the excess rainfall years when seasonal mean standardized rainfall anomaly ≥ 1 and (b, d) for the deficit rainfall years when seasonal mean standardized rainfall anomaly ≤ -1 . Moisture deficit and wind vectors at 850-hPa analysis are illustrated at top panel (a, b) and bottom panel (c, d) shows temperature analysis at 850-hPa. The difference is computed by subtracting long term climatology (1951-2019) from each of the above defined case (a, b and c, d). July-August months are chosen to compute the seasonal mean for each year. The area with gray shading represents the region where the climatology (1951-2019) of surface pressure is less than 850-hPa, and it is masked before the analysis.	56

- 4.5 (a) July-August climatology (negative represents evaporation and positive condensation) and composite anomalies of evaporation (m of water equivalent; scale $\times 10^{-5}$) for (b) excess rainfall years and (c) deficit rainfall years. July-August months are chosen to compute the seasonal mean for each year. The climatology is computed for 1951-2019. The area with gray shading represents the region where the climatology (1951-2019) of surface pressure is less than 850-hPa, and it is masked before the analysis. 57
- 4.6 Seasonal mean (July-August) (a, b) empirical orthogonal function (EOF) with its first two leading modes of mean sea level pressure anomaly and (c, d) linear regression analysis of moisture deficit transport vectors ($\mathbf{V}_{qd}$; $\text{g kg}^{-1} \text{ ms}^{-1} \text{ Pa}^{-1}$) at 850-hPa onto the first two leading principal components i.e. PC1 and PC2. The region with statistically significant ($p < 0.05$) regression coefficients is shown as stippling. The statistical significance is computed using two-tailed t -test (c, d). The analysis is performed for July-August during 1981-2019, and a similar period is chosen for climatology computation to calculate mean sea level pressure anomaly. The region where the climatology (1951-2019) of surface pressure is less than 850-hPa is masked before the analysis. 58
- 4.7 Seasonal mean (July-August) linear regression analysis of (a, b) OLR ($\text{W m}^{-2} \text{ Pa}^{-1}$) and (c, d) precipitation ($\text{mm day}^{-1} \text{ Pa}^{-1}$) onto the first two leading principal components i.e. PC1 and PC2 of mean sea level pressure anomaly. The region with statistically significant ($p < 0.05$) regression coefficients is shown as dashed lines. The statistical significance is computed using two-tailed t -test. The analysis is performed for July-August during 1981-2019, and a similar period is chosen for climatology computation to calculate mean sea level pressure anomaly. The region where the climatology (1951-2019) of surface pressure is less than 850-hPa is masked before the EOF analysis. 59
- 4.8 Seasonal mean (July-August) (a, b) empirical orthogonal function (EOF) with its first two leading modes of mean sea level pressure anomaly and (c, d) linear regression analysis of precipitation ($\text{mm day}^{-1} \text{ Pa}^{-1}$) onto the first two leading principal components i.e. PC1 and PC2. The region with statistically significant ($p < 0.05$) regression coefficients is shown as stippling. The statistical significance is computed using two-tailed t -test (c,d). The analysis is performed for July-August during 1998-2019, and a similar period is chosen for climatology computation to calculate mean sea level pressure anomaly. The TRMM precipitation dataset is used. The region where the climatology (1951-2019) of surface pressure is less than 850-hPa is masked before the EOF analysis. 60

- 4.9 Scatter plot showing the relationship between seasonal mean (July-August) of standardized rainfall anomaly computed over northwestern India and standardized moisture deficit transport vectors (V_{qd}) at 850-hPa calculated over the northern Arabian Sea. The correlation is statistically significant ($p < 0.05$) as revealed by the two-tailed t -test. The analysis is performed for the years 1971-2019, and the 1951-2019 period is chosen for the computation of climatology to construct standardized rainfall anomalies. Northwest India and northern Arabian Sea boxes are indicated in Figs. 4.1 and 4.3, respectively. 61
- 4.10 Seasonal mean (July-August) regression analysis of **(a)** Sea Surface Temperature (SST) anomalies ($^{\circ}\text{C}$ per $^{\circ}\text{C}$) and Precipitation (mm day^{-1} per $^{\circ}\text{C}$) **(b)** IMD and **(c)** GPCP onto the ENSO index. The region with statistically significant ($p < 0.05$) regression coefficients is shown as dashed lines. The statistical significance is computed using two-tailed t -test. The analysis is performed for July-August during 1981-2019 and the 1981-2010 period is chosen for the computation of SST climatology. 63
- 4.11 **(a)** Spatial pattern of the leading mode of EOF (EOF1) of seasonal mean (July-August) meridional wind anomalies at 200-hPa. Linear regression analysis of anomalies of **(b)** meridional wind (m s^{-1}) and **(c)** geopotential height (m) at 200-hPa onto the silk road pattern index (SRPI). The region with statistically significant ($p < 0.05$) regression coefficients is shown as stippling. The statistical significance is computed using two-tailed t -test. The analysis is performed for July-August mean during 1981-2019, and the 1981-2010 period is chosen for the computation of climatology. 64
- 4.12 Seasonal mean (July-August) linear regression analysis of precipitation (mm day^{-1}) for dataset **(a)** IMD and **(b)** GPCP onto the SRPI. The region with statistically significant ($p < 0.05$) regression coefficients is shown as stippling. The statistical significance is computed using two-tailed t -test. The analysis is performed for July-August mean during 1981-2019, and the 1981-2010 period is chosen for the computation of climatology to construct SRPI. The rectangular box shows the NWI domain. 65
- 4.13 Lead-lag composite analysis of anomalies of moisture deficit (g kg^{-1} ; shading) and wind vectors (ms^{-1} ; vectors) at 850-hPa. The analysis is performed for July-August during 1981-2019, and t_0 represents the onset of a break spell identified when daily standardized rainfall anomaly ≤ -1 for atleast three consecutive days. Only statistically significant ($p < 0.05$) values of moisture deficit anomalies are shown. The 1981-2019 period is chosen to compute climatology. 66

4.14	Lead-lag composite analysis of anomalies of moisture deficit (g kg^{-1} ; shading; statistically significant) and wind vectors (ms^{-1} ; vectors) at 850-hPa. The analysis is performed for July-August during 1981-2019, and t_0 represents the onset of an active spell identified when daily standardized rainfall anomaly ≥ 1 for atleast three consecutive days. Only statistically significant ($p < 0.05$) values of moisture deficit anomalies are shown. The 1981-2019 period is chosen to compute climatology.	67
4.15	Lead-lag linear regression of daily moisture deficit transport vectors ($\mathbf{V}_{\text{qd}}$; $\text{kg kg}^{-1} \text{m s}^{-1}$) at 850-hPa onto daily standardized rainfall anomaly computed over the northwestern region (green rectangular box; Fig. 4.1). The statistically significant ($p < 0.05$) regression coefficients are shown as shading. The statistical significance is computed using a two-tailed t -test. The analysis is performed for July-August during 1981-2019, and t_0 represents 01-July. The 1951-2019 period is chosen for the computation of climatology to construct daily standardized rainfall anomalies.	70
5.1	Taylor diagram for model simulated July-August (JA) mean precipitation and zonal winds at 850 hPa. The individual model simulated 1981 – 2014 climatology of JA mean precipitation and zonal winds over ISM domain (40°E – 120°E , 10°S – 30°N) are compared with observations. The precipitation observation is taken from GPCP and winds from ERA5 reanalysis.	77
5.2	Seasonal climatology of precipitation (shading; mm day^{-1}) and wind vectors (vectors; m s^{-1}) at 850-hPa for the July-August months for 14 better performing CMIP6 models as listed in Table 5.2. The precipitation observation is taken from TRMM, and winds from ERA5 reanalysis. The climatology is calculated for the period of 1998-2014.	78
5.3	Taylor diagram for model simulated July-August (JA) mean precipitation and zonal wind at 850 hPa for 14 better performing models. The individual model simulated 1998–2014 climatology of JA mean precipitation and zonal winds over the Indian domain (65°E – 100°E , 5°N – 35°N) are compared with observations. The precipitation observation is taken from TRMM, and zonal winds from ERA5 reanalysis.	80
5.4	Taylor diagram for model simulated July-August (JA) mean precipitation and saturation deficit at 850 hPa. The individual model simulated 1981–2014 climatology of JA mean precipitation and saturation deficit over the ISM domain (40°E – 120°E , 10°S – 30°N) are compared with observations. The precipitation observation is taken from GPCP and the saturation deficit from ERA5 reanalysis.	81

5.5	Meridional wavenumber-frequency spectra of daily precipitation anomaly for 14 better performing CMIP6 models as listed in Table 5.2 and TRMM dataset as observation during JJAS for 15°S-30°N, 60°E-100°E for the year 1998-2014 and the similar period is used to construct climatology	82
5.6	Spatial pattern of EOF (shading; $\times 10^{-2}$) of daily zonal wind anomalies at 850 hPa corresponds to 2 nd leading mode for 14 better performing CMIP6 models as listed in Table 5.2 and ERA5 reanalysis as observation. The EOF analysis is done for the July-August period for the years 1995-2014. The masked orography region is shown with gray shading.	84
5.7	Regression of saturation deficit transport vectors (V_{qd}) on PC2 of daily zonal wind anomalies at 850 hPa for 14 better performing CMIP6 models as listed in Table 5.2 and ERA5 reanalysis dataset as observation. The shading shows statistically significant ($p < 0.05$) regression coefficients ($\text{g kg}^{-1} \text{ms}^{-1}$ per ms^{-1}) of magnitude of saturation deficit transport on PC2. Statistical significance is estimated using a two-tailed t -test. The calculation is done for the July-August period during 1995-2014.	85
5.8	Total number of events of the dry spell, intrusion, and no-intrusion activity for ensemble mean of better 14 CMIP6 models for the period 1995-2014. The error bars indicate the inter-model spread (± 1 standard deviation). Individual model event counts are listed in Table 5.3.	87
5.9	Lead-lag composite of area averaged zonal winds (blue line) and d^2U/dy^2 (black line) for the dry air intrusion events for observation. A three-day running mean is applied to lead-lag time series. The lag-0 corresponds to the first day of intrusion. The area average is done over 55°E - 75°E and 10°N - 20°N. The analysis period is 1995-2014 and the ERA5 reanalysis dataset is used as observation. The onset index of dry air intrusion is listed in Table 5.4.	89
5.10	Lead-lag composite of area averaged zonal winds (blue line) and d^2U/dy^2 (black line) for the dry air intrusion events for ensemble mean of 14 better performing CMIP6 models as listed in Table 5.2. A three-day running mean is applied to lead-lag time series. The lag-0 corresponds to the first day of intrusion. The area average is done over 55°E - 75°E and 10°N - 20°N. The analysis period is 1995-2014. The onset index of dry air intrusion is listed in Table 5.4.	90

5.11	Lead-lag composites of time evolution of vertical structure of equivalent potential temperature anomalies (K) (a) Observation (ERA5) and (b) Ensemble mean of 14 better performing CMIP6 models (Table 5.2) averaged over northwestern India (65°E–75°E, 20°N–30°N) for dry air intrusion events (shading) and for monsoon breaks without dry air intrusion (contour; dashed contour shows negative value and contour value ranges from -2 to 2 with a level spacing of 0.4). Lag 0 corresponds to the day of maximum of dry air intrusion and no-intrusion events. The calculation is done for the July-August period during 1995-2014.	92
5.12	Lead-lag composites of daily anomalies of specific humidity (g kg^{-1}) at 316K isentropic level for all dry air intrusion events in July–August during 1995-2014 for observation. t_0 corresponds to the day of maximum of dry air intrusion. Stippling represents statistically significant ($p < 0.05$) regions. ERA5 reanalysis dataset is used as observation.	93
5.13	Lead-lag composites of daily specific humidity anomalies (g kg^{-1}) at 316K isentropic level for all dry air intrusion events in July–August during 1995-2014 for ensemble mean of 14 better performing CMIP6 models as listed in Table 5.2. t_0 corresponds to the day of maximum of dry air intrusion.	94
5.14	Lead-lag composites of daily anomalies of relative vorticity (shading; $\times 10^{-5}\text{s}^{-1}$) and wind vectors (ms^{-1}) at 316K isentropic level for all dry air intrusion events in July–August during 1995-2014 for observation. t_0 corresponds to the day of maximum of dry air intrusion. Stippling represents the region where anomalies of relative vorticity is statistically significant ($p < 0.05$). ERA5 reanalysis dataset is used as observation.	95
5.15	Lead-lag composites of daily anomalies of relative vorticity (shading; $\times 10^{-5}\text{s}^{-1}$) and wind vectors (ms^{-1}) at 316K isentropic level for all dry air intrusion events in July–August during 1995-2014 for ensemble mean of 14 better performing CMIP6 models as listed in Table 5.2. t_0 corresponds to the day of maximum of dry air intrusion.	97
5.16	Lead-lag composites of daily equivalent potential temperature anomalies (K) at 700 hPa for all dry air intrusion events in July–August during 1995-2014 for ensemble mean of 14 better performing CMIP6 models as listed in Table 5.2. t_0 corresponds to the day of maximum of dry air intrusion.	98
5.17	Lead-lag composites of daily anomalies of equivalent potential temperature (K) at 700-hPa for all dry air intrusion events in July–August during 1995-2014 for observation. t_0 corresponds to the day of maximum of dry air intrusion. Stippling represents statistically significant ($p < 0.05$) regions. ERA5 reanalysis dataset is used as observation.	99

5.18	Composite of equivalent potential temperature at 850 hPa (shading; K) (a) Dry spells with intrusion, (b) Dry spells with no-intrusion and (c) Active spells in July–August during 1995–2014 for ensemble mean of 14 better performing CMIP6 models as listed in Table 5.2.	100
5.19	Climatology of CTRL experiments are shown. Simulated July–August mean of precipitation (mm day^{-1}) (left) and sea level pressure (hPa) (right) are shown respectively for the simulated ten years.	102
5.20	Climatology of CTRL experiments are shown. Simulated July–August mean of wind vectors (vectors; m s^{-1}) (left), saturation deficit (shading; g kg^{-1}) (left) and temperature (shading; K) (right) at 850 hPa are shown respectively for the simulated ten years.	103
5.21	Seasonal mean difference from CTRL experiments are shown. Simulated July–August mean of precipitation (mm day^{-1}) (left) and sea level pressure (hPa) (right) are shown respectively for the two sensitivity experiments for the simulated ten years.	104
5.22	Seasonal mean difference from CTRL experiments are shown. Simulated July–August mean of wind vectors (vectors; m s^{-1}) (left), saturation deficit (shading; g kg^{-1}) (left) and temperature (shading; K) (right) at 850 hPa are shown respectively for the two sensitivity experiments for the simulated ten years.	105
5.23	Seasonal mean (July–August) of precipitation (mm day^{-1}) for 15 CMIP6 models for historical experiments. The period of analysis is 1985–2009.	106
5.24	Taylor diagram for model simulated July–August (JA) mean precipitation and saturation deficit at 850 hPa. The individual model simulated 1985–2009 climatology of JA mean precipitation and saturation deficit over ISM domain (40°E – 120°E , 10°S – 30°N) are compared with observations. The precipitation observation is taken from GPCP and the saturation deficit from ERA5 reanalysis.	107
5.25	Seasonal mean (July–August) of moisture deficit (shading; g kg^{-1}) and wind vectors (vectors; m s^{-1}) at 850 hPa for 15 CMIP6 models for historical experiments. The period of analysis is 1985–2009.	108
5.26	Seasonal mean (July–August) difference (SSP 585 – hist.) of precipitation (mm day^{-1}) for 15 CMIP6 models. The period of the analysis is 1985–2009 (hist.) and 2075–2099 (SSP 585).	109
5.27	Seasonal mean (July–August) difference (SSP 585 – hist.) of moisture deficit (shading; g kg^{-1}) and wind vectors (vectors; m s^{-1}) at 850 hPa for 15 CMIP6. The period of the analysis is 1985–2009 (hist.) and 2075–2099 (SSP 585).	110
5.28	Ensemble mean difference (SSP 585 – hist.) of moisture deficit (shading; g kg^{-1}) and wind vectors (vectors; m s^{-1}) at 850 hPa for the CMIP6 models (a) break days increases and (b) break days decreases. The period of the analysis is 1985–2009 (hist.) and 2075–2099 (SSP 585).	111

5.29 Ensemble mean difference (SSP 585 - hist.) of precipitation (mm day ⁻¹) for the CMIP6 models (a) break days increases and (b) break days decreases. The period of the analysis is 1985–2009 (hist.) and 2075–2099 (SSP 585).	112
---	-----

List of Tables

1.1	Thesis structure	13
3.1	The start and end dates of dry air intrusion events during monsoon break spells in 1981-2014 period	33
3.2	The dates of monsoon break spells without a dry air intrusion	33
4.1	Excess rainfall years and the status of ENSO	68
4.2	Deficit rainfall years and the status of ENSO	68
5.1	List of 31 CMIP6 global climate models used in this study	83
5.2	Listing the pattern correlation (PC), normalized standard deviation (Ratio), and Taylor skill score (TSS) with respect to observation for precipitation (saturation deficit at 850-hPa) of all CMIP6 climate models used to create the Taylor diagram in Fig. 5.1 and 5.4.	86
5.3	Total number of events (total duration in days) for a dry spell, intrusion and no-intrusion activity for each model (14 better performing models and observation). The analysis year is 1995-2014 for July-August.	88
5.4	The index of the onset of dry air intrusion for each event is listed for 14 individual models and observation during the ISM break phase. The analysis covered the time period of July and August from 1995 to 2014. The index starts from July 1, 1995, and ends on August 31, 2014, with the start index being 1 and the end index being 1240.	91
5.5	The deviation in total number of break days is computed for 14 CMIP6 models. The deviation is the difference between SSP 585 and historical experiments. The period of analysis is 1985-2009 (historical) and 2075-2099 (SSP 585) for July and August.	113

Abbreviations

AGCM	Atmospheric General Circulation Model
BoB	Bay of Bengal
CAM	Community Atmosphere Model
CESM	Community Earth System Model
CMIP6	Coupled Model Intercomparison Project Phase Sixth
DAI	Dry Air Intrusion Index
ECMWF	European Centre for Medium-Range Weather Forecasts
ENSO	El Niño Southern Oscillation
IOD	Indian Ocean Dipole
IMD	India Meteorological Department
ISM	Indian Summer Monsoon
ISMR	Indian Summer Monsoon Rainfall
LLJ	Low Level Jet
NOAA	National Oceanographic and Atmospheric Administration
NWI	Northwestern (or Northwest) India
OLR	Outgoing Longwave Radiation
SENS	Sensitivity
SRP	Silk Road Pattern
SST	Sea Surface Temperature
TRMM	Tropical Rainfall Measuring Mission